

Editorial

Decadal Changes of Earth's Outgoing Longwave Radiation

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Abstract: The Earth Radiation Budget (ERB) at the top of the atmosphere quantifies how the earth gains energy from the sun and loses energy to space. Its monitoring is of fundamental importance for understanding ongoing climate change. In this paper, decadal changes of the Outgoing Longwave Radiation (OLR) as measured by the Clouds and Earth's Radiant Energy System from 2000 to 2018, the Earth Radiation Budget Experiment from 1985 to 1998, and the High-resolution Infrared Radiation Sounder from 1985 to 2018 are analysed. The OLR has been rising since 1985, and correlates well with the rising global temperature. An observational estimate of the derivative of the OLR with respect to temperature of $2.93 \pm 0.3 \text{ W/m}^2\text{K}$ is obtained. The regional patterns of the observed OLR change from 1985–2000 to 2001–2017 show a warming pattern in the Northern Hemisphere in particular in the Arctic, as well as tropical cloudiness changes related to a strengthening of La Niña.

Keywords: earth radiation budget; outgoing longwave radiation

1. Introduction

Earth's climate is determined by the Earth Radiation Budget (ERB). For a climate in equilibrium, the gain of energy from Absorbed Solar Radiation (ASR) is in balance with the loss of energy through Outgoing Longwave Radiation (OLR). Any perturbation of this radiative energy balance at the Top of Earth's Atmosphere (TOA) is known as a radiative forcing, and is a driver of climate change [1]. Compared to the pre-industrial period, the increase of Greenhouse Gases (GHG) primarily CO_2 reduces the OLR; this represents a positive (heating) radiative forcing [2]. This positive GHG radiative forcing is partially compensated by a negative (cooling) radiative forcing of aerosols, which increase the amount of Reflected Solar Radiation (RSR) and hence decrease the ASR [3]. Variations of the incoming solar radiation could also create a radiative forcing, but in practice the solar irradiance variations are small [4].

For understanding current and past climate change, and for increasing the confidence with which future climate change can be predicted, the analysis of long-term changes of the ERB is essential. The measurement of the ERB from space has a long history [5], but the availability of long-term homogeneous ERB Climate Data Records (CDRs) is relatively recent.

In this paper, we analyse the long-term variation of selected OLR CDR's in order to increase our understanding of climate change. In Section 2, we describe the used data and the main features of the climatological mean OLR. In Section 3, we analyse the decadal change of the global annual mean OLR over the period 1985–2018. In Section 4, we investigate the regional patterns of the OLR change. In Section 5, we discuss our results.

2. Outgoing Longwave Radiation

The earth surface and the constituents of the earth's atmosphere emit thermal radiation. The OLR is defined as the emitted thermal radiation that is radiated to space at the TOA. In this paper, we consider the following long-term measurements of the OLR:

- Broadband scanner measurements made by the National Aeronautics and Space Administration (NASA) Clouds and the Earth's Radiant Energy System (CERES) [6] instruments. We use the CERES EBAF Ed4.0 following [7] from 2000 to 2018.
- Broadband non scanner measurements made with the NASA Earth Radiation Budget Experiment (ERBE) [8] Earth Radiation Budget Satellite (ERBS) Wide Field Of View (WFOV) instrument. We use the ERBE WFOV Ed4.0 following [9] from 1985 to 1998.
- Broadband estimates derived from multichannel narrowband measurements made by the High-resolution Infrared Radiation Sounder (HIRS) instruments on the National Oceanic and Atmospheric Administration (NOAA) and Eumetsat's Metop satellites. We use the HIRS OLR version 2.7 following [10]. Intersatellite calibrations were performed to remove possible discontinuity in the HIRS OLR product from different satellites. In the 2012, Gewex Radiative Flux Assessment (RFA) [11], the HIRS OLR was the only long-term OLR dataset comparing favourably with both the ERBE scanner and CERES reference datasets for the periods of 1985–1989 and 2001–2004, respectively.

Following [5], we put the HIRS and CERES OLR on the same absolute scale. We make the OLR equal to the reference value of 238 W/m^2 for the period 2005–2014. Therefore, we multiply the HIRS OLR with a factor 0.997134 and we multiply the CERES OLR with a factor 0.990837. The ERBE WFOV OLR is made equal to the HIRS OLR over the period 1985–1998. Therefore, we multiply the ERBE WFOV OLR by 0.995169.

Figure 1 shows the 17-year CERES mean 'all sky' OLR over the period 2001–2017. Figure 2 shows the corresponding clear sky OLR, and Figure 3 shows the OLR cloud effect defined as the all sky OLR minus the clear sky OLR.

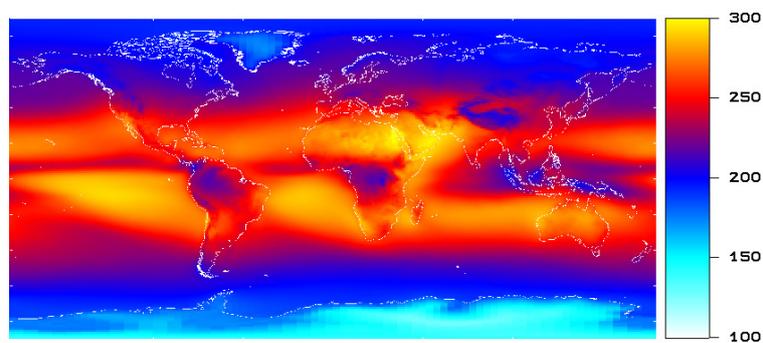


Figure 1. 2001–2017 mean all-sky CERES OLR. The scale is in W/m^2 .

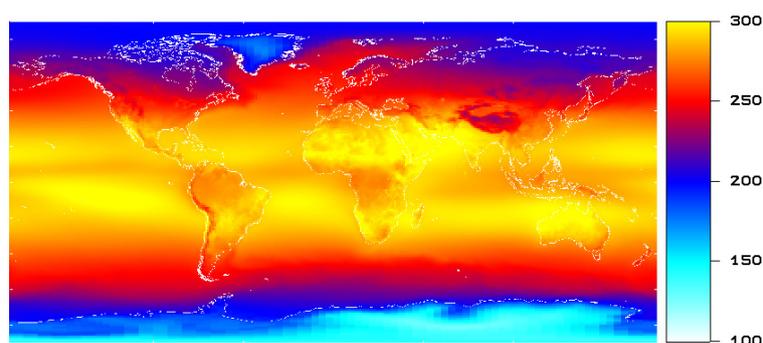


Figure 2. 2001–2017 mean clear-sky CERES OLR. The scale is in W/m^2 .

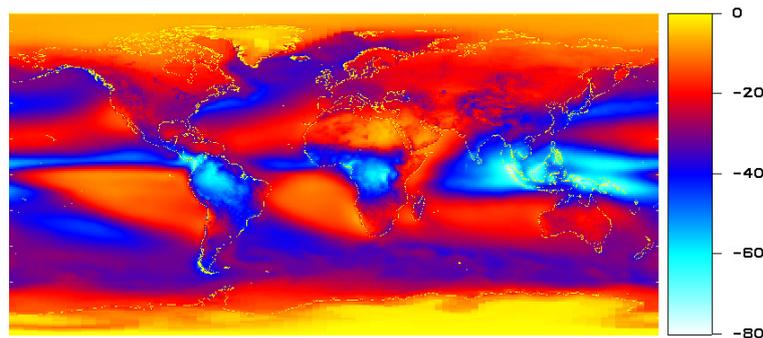


Figure 3. 2001–2017 mean cloud effect CERES OLR. The scale is in W/m^2 .

The clear sky OLR (Figure 2) shows the variation of high thermal emission (yellow and red colors) linked to high surface temperatures at the equator towards low thermal emission (blue and light blue colors) linked to low surface temperatures at the poles. In the equatorial convective regions, the clear sky thermal emission is slightly reduced (light red color) due to water vapour absorption compared to the subtropical high pressure regions (yellow color).

Near the equator, tropical convective clouds cause OLR minima that can be seen in Figures 1 (blue color) and 3 (light blue and blue color). Tropical convection is caused by the diurnal heating of the earth surface by the sun [12]. This heating causes a faster temperature rise over land than over ocean, hence tropical convection, and the associated OLR minima, form preferentially over land.

At midlatitudes, the OLR is also reduced by clouds—particularly over the oceans—as can be seen in Figure 3 (blue color). In the Northern Hemisphere, OLR minima form over ocean, and correspond more or less to the Icelandic and Aleutian low pressure systems [13]. These low pressure systems are the strongest in winter, when the ocean is warmer than the land. In the Southern Hemisphere, there exists a zonally extended low pressure system and corresponding OLR minimum over the Southern Ocean.

3. Global Annual Mean Outgoing Longwave Radiation Change 1985–2017

Figure 4 shows the running yearly mean of the global average OLR as measured by HIRS, ERBE and CERES, as well as a linear regression model based on the yearly NASA GISS T anomaly [14] over the period 1985–2018.

The linear regression model is

$$OLR(\Delta T) = OLR_0 + (dOLT/dT) * \Delta T \quad (1)$$

with ΔT the NASA GISS T anomaly, and OLR_0 and $dOLT/dT$ as regression parameters.

The Root Mean Square Difference (RMSD) between the data plotted in Figure 4 give an indication of their accuracy. For the measurements, the RMSD between HIRS and CERES is $0.15 W/m^2$, and the RMSD between HIRS and ERBE is $0.53 W/m^2$. For the regression model, the RMSD between the model and HIRS is $0.34 W/m^2$.

The slope $dOLT/dT$ from Equation (1) is the derivative of the OLR with respect to global surface temperature. It is the longwave part of the climate feedback parameter as defined in [15]. The measured value with 1σ uncertainty interval we obtain is

$$dOLT/dT = 2.93 + / - 0.3 W/m^2K. \quad (2)$$

It can be expected that the OLR rises with increasing global temperature. This is a mechanism by which earth can return to a climate equilibrium after an initial perturbation by an external radiative forcing. Our measured value of $dOLT/dT$ can be compared with the model values from [15]. If earth were a blackbody without an atmosphere, the derivation of the Planck curve yields

$dOLR/dT = 3.3 \text{ W/m}^2\text{K}$. Inclusion of the water vapour feedback—if earth warms, its atmosphere contains more water vapour, which strengthens the greenhouse effect and yields a ‘clear-sky’ $dOLR/dT = 2.2 \text{ W/m}^2\text{K}$. $dOLR/dT$ is determined for 19 General Circulation Models (GCMs) in Table 10 of [15]. The model ‘all-sky’ $dOLR/dT$ falls in the range 2.04–3.13. From the 19 models of $dOLR/dT$, only the six highest ones fall in the interval of Equation (2). Comparing the measured ‘all-sky’ $dOLR/dT$ of Equation (2) with the ‘clear-sky’ $dOLR/dT$ of $2.2 \text{ W/m}^2\text{K}$, we can conclude there exists a ‘longwave cloud thinning effect’: as the earth warms, it contains less clouds, and becomes a more effective radiator. This cloud thinning effect is underestimated in most of the models.

After the 1991 Pinatubo eruption, there was a temporary drop in the OLR measured by ERBE and by HIRS, as well as in the global temperature. This episode is discussed in detail in [16].

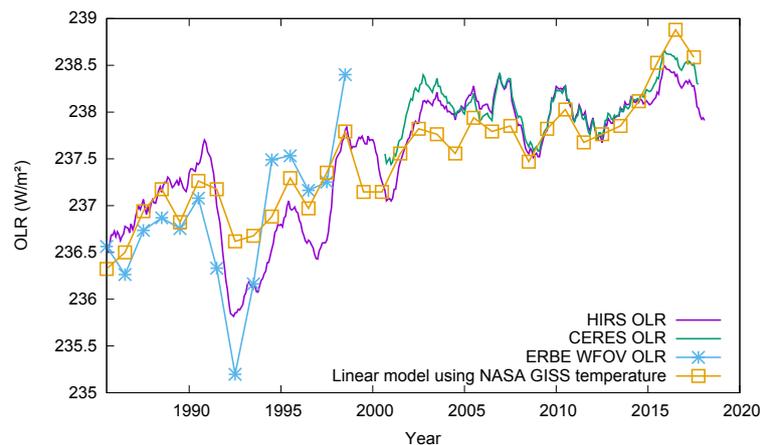


Figure 4. Purple curve: HIRS OLR. Green curve: CERES OLR. Blue curve: ERBE WFOV OLR. Orange curve: linear regression model based on NASA GISS T anomaly.

4. Regional Annual Mean Outgoing Longwave Radiation Change 1985–2017

Figure 5 shows the mean HIRS OLR over the period 2001–2017 minus the mean over the period 1985–2000.

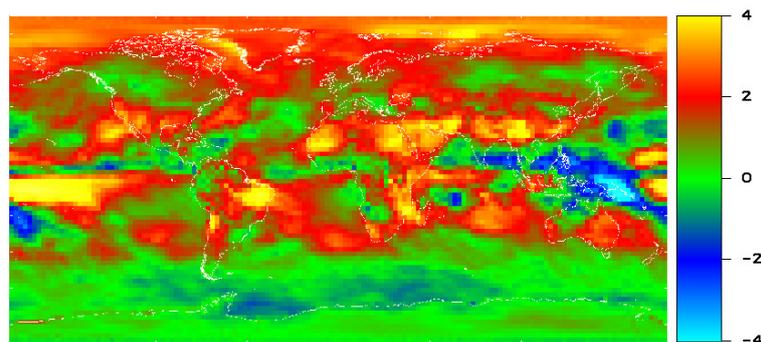


Figure 5. Change of HIRS OLR from 1985–2000 to 2001–2017. The scale is in W/m^2 .

In Figure 5, we can see an increase of the OLR (red-yellow colors) in the subtropical high pressure areas, part of the North-Hemisphere mid-latitude regions, and the Arctic, where the temperature rise is the highest. For reference, Figure 6 shows the NASA GISS temperature change from 1985–2000 to 2001–2017 [14].

In the tropics, changes in OLR are visible in Figure 5, which are related to El Niño/La Niña variability.

We use the Multivariate El Niño Index (MEI) from [17] to characterise the El Niño/La Niña conditions. The MEI is a dimensionless index where a value of 1 indicates a medium strength El Niño

and a value of -1 indicates a medium strength La Niña. From the monthly values provided by [17], we produce running yearly means.

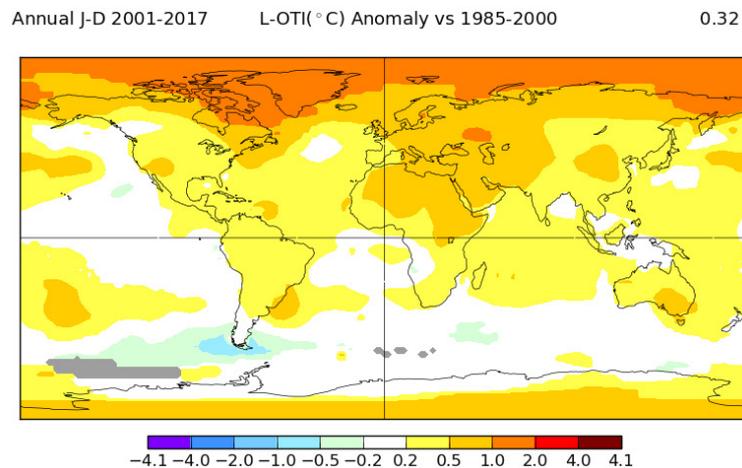


Figure 6. NASA GISS T change from 1985–2000 to 2001–2017.

Figure 7 shows the variability of the HIRS OLR between El Niño years identified by a positive MEI and La Niña years identified by a negative MEI. Regions with a decreased OLR (blue colors in Figure 7) have an increased precipitation during La Niña, regions with an increased OLR (red colors in Figure 7) have a decreased precipitation during La Niña. The tropical spatial pattern of Figure 5 shows some resemblance with the tropical spatial pattern of Figure 7. This indicates a relative strengthening of La Niña after 2000. This is consistent with our earlier findings [5].

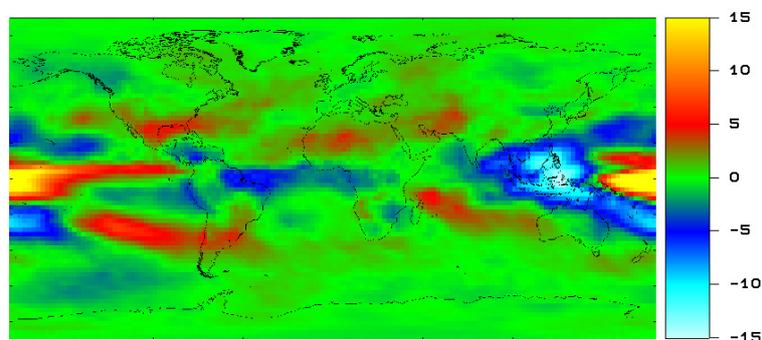


Figure 7. Change of HIRS OLR between El Niño years and La Niña years. The scale is in W/m^2 .

For a more detailed investigation of the decadal ENSO changes, in Figure 8, we show the yearly MEI index $MEI(y)$ together with the Cumulative MEI index $CMEI(y)$ defined as

$$CMEI(y) = \sum_{1950 \leq y' \leq y} MEI(y'). \quad (3)$$

From 1950 to 1975, $CMEI(y)$ was decreasing, indicating a period of strong La Niñas. From 1975 to 1998, $CMEI(y)$ was increasing, indicating a period of strong El Niños. From 1998 to 2014, $CMEI(y)$ was relatively flat, indicating equal strength between La Niña and El Niño during the period. Compared to the period before 1998, this means that there was a relative strengthening of La Niña, and this explains the La Niña pattern we have found above in the tropical OLR variations in Figure 5.

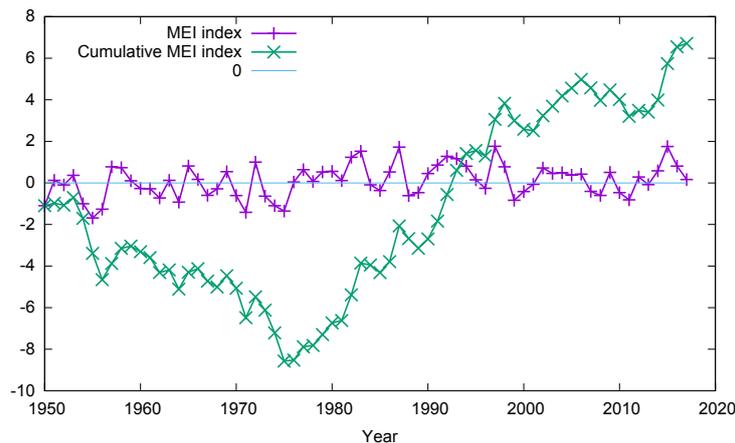


Figure 8. Purple: Multivariate El Niño Index (MEI); green: cumulative MEI index.

5. Discussion

We have analysed the long-term OLR change measured by the HIRS instruments over the period 1985–2017. The HIRS measurements were previously found to be of good quality [10,11] and are further validated in this study by comparison with the CERES Edition 4 OLR after 2000, and by comparison with the ERBE Edition 4 OLR before 1998. The RMSD with CERES is 0.15 W/m^2 , and the RMSD with ERBE is 0.53 W/m^2 .

From the joint analysis of the HIRS OLR and the NASA GISS global surface temperature anomaly, we derive an empirical estimate of the longwave climate feedback parameter $dOLR/dT$ of $2.93 \pm 0.3 \text{ W/m}^2\text{K}$ 1σ uncertainty interval. It is well known [15,18] that the ability of Global Climate Models (GCMs) to predict future climate change is limited by the knowledge of cloud feedback processes. Our empirical measurement of $dOLR/dT$ allows for putting a constraint on the GCM longwave cloud feedback processes. Most of the GCMs in the [15] study have too low of a $dOLR/dT$, indicating an underestimating of the ‘longwave cloud thinning effect’: as the earth warms, it contains less clouds and becomes a more effective radiator.

The spatial analysis of the regional OLR change from 1985–2000 to 2001–2017 shows a ‘clear-sky effect’ mixed a ‘tropical cloud’ effect.

Concerning the clear sky effect, we see regional changes of the OLR which are correlated with surface temperature changes. In the Arctic—where the strongest temperature increase occurs—we also see a strong increase in the OLR. In general, in the Northern Hemisphere—where the surface temperature increase is stronger than in the Southern Hemisphere—we also see an OLR increase.

Concerning the tropical cloud effect, we see regional patterns in the changes of the OLR, which are suggesting a relative strengthening of La Niña conditions compared to El Niño conditions. These changes imply societally important regional changes in precipitation. The relative La Niña strengthening can also be seen in the ‘cumulative MEI index’ that we have introduced.

Due to its fundamental importance for climate change, the monitoring and analysis of Earth’s Outgoing Longwave Radiation should be continued.

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